Tracer Tests with Resistivity Detection to Estimate Seepage Velocity in the Biscayne Aquifer, Florida

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Water Management in the Everglades Restoration



NATURAL SYSTEM (circa 1850)



MANAGED SYSTEM (1995 Landscape)



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Seepage Management for Everglades Restoration



- Maintain and increase groundwater levels to the west
- Improve hydroperiod in Everglades NP
- Keep high quality surface/ground water in ENP
- Maintain flood control protection as CERP projects increase surface and groundwater flows into ENP
- Seepage management using cut-off wall



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Hydrogeologic Conditions at the L-30 Test Site

Natural Surface Water Flow

Tracer Test Wellfield

30 Leve

Tamiami Trail Culverts

Northeast Shark River Slough

Everglades National Park

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Hydrogeologic Conditions at the L-30 Test Site

Regional Groundwater Flow Direction

120 to 140 ° azimuth Estimated 162,000 acre-ft seepage loss

Ezone Google

Tracer Test

Wellfield

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Hydrogeologic Characteristics of the Biscayne Aquifer

- Marine limestone karstic aquifer material
- Heterogeneous pore network
- Permeability values range over 13 orders of magnitude (Cunningham et al., 2011)
- Drinking water supply for 2.5M in Miami-Dade Co.





Photos from Cunningham and Sukop (2011) USGS OFR-1037



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Hydrogeologic Conditions at the L-30 Test Site

Fort Thompson Formation includes the Biscayne Aquifer



Optical Borehole Images in Wellfield Boreholes



CROSS-SECTION ACROSS "WINDOW"

Instrumented Wellfield at the L-30 Test Site



Photograph courtesy Jim Brock, Rapid Creek Research





- Wells have 2-ft screen intervals at upper flow zones
 Hourly measurement of temperature, specific conductance, water level, flow velocity and flow direction
- Monitoring effort will compare wet and dry season conditions
- Resistivity transects superposed on the well field for test



Quantifying Groundwater Flow Velocity

Approaches

 Numerical Groundwater Flow Model Calculate Darcian (non-Darcian?) flow velocity Heat-Pulse Borehole Flow Meters Do FM velocities represent aquifer flows? Tracer Tests using Bromide Tracer Detecting breakthrough in down-gradient wells. Needle in a haystack? Resistivity survey provides better spatial resolution of ionic tracer transport Flow induced from dense saline tracer?

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Electrical Resistivity



• Electrical Current injected through two current electrodes.

 Voltage drop is measured across potential electrodes.

• Electrode array is expanded to increase depth of penetration.

 Resistivity of formation/fluids measured in ohm-meters (Ω-m).

• Modern systems use many electrodes with automated switching.



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Resistivity Survey Methods and Advantages

- Ionic Tracer (KBr) to provide resistivity contrast with fresh groundwater.
- Resistivity survey consists of 4 lines parallel to L-30, 1 line perpendicular. Tracer detected in X-Y-Z axes
- Each line is 270 ft long with electrodes every 10 ft.
- Resistivity contrast detected at full depth of Biscayne Aquifer (65 ft below land surface)
- Can be conducted in wet or dry conditions
- Rapid data collection: 4 lines surveyed in 1 hour
- Results can be viewed in field, so tracer detection frequency can be evaluated and modified



Field Work: May 2010 (wet) and April 2011 (dry)



May 2010 Test: 50 lbs KBr in 100 gal potable water with 100 gal potable water chaser

April 2011 Test: 500 lbs KBr in 500 gal potable water, with 500 gal potable water chaser

Dosing well open interval restricted with packer. Open interval at -0.89 to -12.09 ft NGVD





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May 2010 Tracer Test Background Conditions: Line 2, Time 0 Before Tracer Dosing



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Results of the "Wet" Season Test (May 2010)

- Resistivity changes are subtle from this tracer test
- Resistivity reduction was observed only in Line 2 (10 ft from dosing well), 54 min after dosing.
 Groundwater flow rate: 10 ft/0.9 hr = 266 ft/day
 - More subtle response appeared in Line 3 (35 ft from dosing well), 365 min after dosing.
 Groundwater flow rate: 35 ft/6.1 hr = 138 ft/day
- Test results suggest that the next experiment should have a larger volume dose of KBr tracer



April 2011 Tracer Test Background Conditions: Line 2, Time 0 Before Tracer Dosing



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April 2011 Tracer Test: Line 2, time 0 to 6 hours **INCREASING** APPROXIMATE LOCATION OF MW-7 WITH INJECTION TIME BETWEEN ABOUT 4 AND 18' BELOW GRADE VIIIHI. 210 170 150 240270 +5.0200 Elevation (Feet NGVD29) -19.9 155 -44.8 110 0 HR -69.7 65 -94.7 20.0 Iteration = 2 RMS = 2.34% L2 = 0.61 Electrode Spacing = 10 ft Ohm-m 30 60 120 150 210 240 270 +5.0200 Elevation (Feet NGVD29) -19.9 155 3.1 HR -44.8 110 -69.7 65 INITIAL TRACER RESPONSE -94.7 20.0 Iteration = 2 RMS = 2.80% L2 = 0.87 Electrode Spacing = 10 ft Ohm-m 120 150 210 240 270 +5.0Elevation (Feet NGVD29) -17.2 155 5.1 HR -39.4 110 -61.6 65 -83.8 20.0 heration = 3 RMS = 9.10% 1.2 = 9.19 Electrode Spacing = 10 ft Ohm-m 120 150 240 270 +5.0200 Elevation (Feet NGVD29) -17.2 155 6.0 HR -39.4 · 110 -61.6 65 SLIGHT REDUCTION -83.8 20.0 Iteration = 3 RMS = 8.90% L2 = 8.79 Electrod: Spacing = 10 ft

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Results of the "Dry" Season Test (April 2011)

- Resistivity changes are clearer in this tracer test due to increased frequency of resistivity measurements
- Maximum resistivity reduction was observed in Line 2 (10 ft from dosing well), 5.1 hr after dosing.
 Groundwater flow rate: 10 ft/5.1 hr = 47 ft/day
- More subtle response appeared in Line 3 (35 ft from dosing well), 5.3 hr after dosing.
 - Groundwater flow rate: 35 ft/5.1 hr = 158 ft/day
- Some density effects observed. Tracer detected at 2.9 hours in upgradient Line 1.



How significant are density effects?

Archie's Equation relates electrical resistivity to porosity and brine saturation

Archie's Equator: $R_W = \emptyset^M R_O$ $R_W = .3^2 R_O$ $R_W = .09 R_O$

 \emptyset = porosity = 20 - 70% use 0.3 for Biscayne

 $M = cementation \ factor \approx 2$ Ranges from $\approx 1.3 \ to \ 2.0$

Ro = Saturated Formation Resistivity

					K ₃ /M ³ *** Density at
Line	R _{O (min)}	$R_W(\Omega m)$	S.C. (μS/cm)	ppm**	25°C
1	65	5.9	1,700	830	997.7
2	80	7.2	1,400	690	997.6
3	105	9.5	1,100	540	997.5
5	112	10	1,000	490	997.4
MW-8	720 µS/cm	14	720	350	997.3
MW-6	580 μS/cm	17	580	280	997.3
Seawater				30,000	1,019.6

 $1 \,\mu\text{S/cm} = 10,000 \,\,\Omega\text{m} = 0.0001 \,\,\text{mho/m}$

Results and Interpretations

 April 2011 test data is interpreted as the best representation of solute flow through the Biscayne Aquifer. Groundwater flow rates are estimated as:

Line 2 is 48 ft/day

Line 3 is 158 ft/day

- Multiple 2-D resistivity transects enable 3-D definition of tracer distribution and transport
- Maximum depth of tracer approximately -35 ft NGVD29



Questions ?

For more information see http://www.evergladesplan.org

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